

U.S. Department of the Interior
U.S. Geological Survey

SIMULATION OF THE EFFECTS OF DEVELOPMENT OF THE GROUND-WATER FLOW SYSTEM OF LONG ISLAND, NEW YORK

Water-Resources Investigations Report 98-4069

Prepared in cooperation with the
NASSAU COUNTY DEPARTMENT OF PUBLIC WORKS,
SUFFOLK COUNTY DEPARTMENT OF HEALTH SERVICES,
SUFFOLK COUNTY WATER AUTHORITY, and the
NEW YORK CITY DEPARTMENT OF ENVIRONMENTAL PROTECTION

The distribution of ground-water pumping as represented in the model is shown by county and model layer (aquifer) in table 12. The greatest increase in pumping is expected to be from the basal zone of the Magothy aquifer (layer 3) in Suffolk County, and the highest rate of pumping will remain in Nassau County from the basal zone of the Magothy aquifer. Projected changes from the distribution of pumping in 1983 to 2020 can be evaluated through comparison of table 10 and table 12.

Table 12. Distribution of ground-water pumping for public-supply, industrial-commercial, and agricultural uses for the year 2020, as represented in model

County	Model layer				Total
	1 (water table)	2 (Magothy and Jameco)	3 Lloyd	4	
Kings and Queens	24	2	17	3	46
Nassau	10	10	181	13	214
Western Suffolk	26	2	79	0	107
Eastern Suffolk	37	15	39	0	91
Total	97	29	316	16	458

Ground-Water System Response

The predicted hydrologic response to the water-supply development strategy for the year 2020 is presented in terms of (1) base flow; (2) movement of the saltwater-freshwater interface; (3) ground-water levels and flow patterns; and (4) the ground-water budget. These results provide a guide to water-resource managers who must define acceptable levels for the adverse effects of development and modify development strategies to meet these levels. The predicted response also is compared with simulated results for the predevelopment and recent hydrologic conditions to demonstrate the evolution of the development of the Long Island ground-water system.

Base Flow

Model predictions of base flow of major streams for the 2020 water-supply strategy are presented in table 13. Streams are represented in the model as drains, similar to simulations of the 1960's drought, however, recent stressed conditions were used as a baseline. Recovery of ground-water levels in Queens and western Nassau County is expected to increase base flow and restore flow in some dry stream channels (table 13 and fig. 31); the base flow of Flushing Creek, Springfield Stream, Simonsons Stream, Valley Stream, Motts Creek and Pines Brook will increase. From South Pond in western Nassau County eastward, however, base flow in all streams will decrease.

Base flow in East Meadow Brook, Bellmore Creek, and Massapequa Creek in Nassau County are estimated to decrease the most—their combined flows will decrease 92 percent, from 22.3 ft³/s during 1968-83 to 1.8 ft³/s by the year 2020. East Meadow Brook is projected to be dry from its headwaters to the gage. Base flow of Santapogue Creek, Carlls River, and Sampawams Creek in western Suffolk County together will decrease to about 60 percent of their flow during 1968-83. As indicated in the analysis of the 1960's drought, long streams that extend far inland are affected most severely; this is evident from comparing the estimated base flow in Massapequa Creek, Bellmore Creek, and East Meadow Brook with Milburn Creek, Cedar Swamp Creek, and Carman Creek in table 13. Streams east of Nissequogue and Connetquot Rivers will be less severely affected than those to the west because the increase in stress will be smaller and because the effects of stress in the west do not propagate past these large streams.

Saltwater-Freshwater Interface

The movement of the saltwater-freshwater interface between 1983 and 2020 cannot be determined by the islandwide model. Movement of the interface was assumed to be

Table 13. Average base flow of major streams on Long Island, estimated for predevelopment and during 1968-83 and predicted for the year 2020

Map number (fig. 3)	Stream name	Period			Map number (fig. 3)	Stream name	Period		
		Predevelopment	1968-83	2020			Predevelopment	1968-83	2020
1	Jamaica Creek	17.9	0.0	0.0	17	Sampawams Creek	9.9	6.7	3.6
2	Springfield Stream	7.9	0.0	0.1	18	Penataquit Creek	6.8	6.5	5.0
3	Simonsons (Brookfield) Stream	9.6	0.3	2.9	19	Pardees and Orowoc Creeks	10.3	8.9	6.9
4	Valley Stream	14.3	0.3	1.7	20	Rattlesnake Brook	9.2	8.8	8.5
5	Motts Creek	6.4	2.1	4.3	21	Connetquot River	36.0	34.6	31.0
6	Pines Brook	13.0	0.5	1.0	22	Green Creek	6.5	6.5*	6.5
7	South Pond	20.0	0.4	0.1	23	Patchogue River	18.9	18.9*	18.4
8	Parsonage Creek	8.1	4.5	3.9	24	Swan River	13.3	13.3*	13.0
9	Milburn Creek	13.0	6.9	4.3	25	Carmans River	24.9	24.9*	24.1
10	East Meadow Brook	15.3	6.3	0.0	26	Forge River	9.6	9.6*	9.1
11	Cedar Swamp Creek	9.5	6.8	2.8	27	Little River	7.4	7.4*	7.4
12	Bellmore Creek	14.6	9.4	1.5	28	Peconic River	37.4	37.4*	35.7
13	Massapequa Creek	* 12.0	6.6	0.3	29	Nissequogue River	41.7	40.2	37.1
14	Carman Creek	6.8	6.7	2.6	30	Mill Neck Creek	7.0	5.6	3.1
15	Santapogue Creek	10.0	8.0	4.9	31	Glen Cove Creek	8.7	3.7	1.8
16	Carlls River	27.3	20.5	11.9	32	Flushing Creek	21.5	7.8	15.5

* Assumed to be the same as under predevelopment conditions because development is minimal, and records indicate no decrease in base flow from predevelopment conditions.

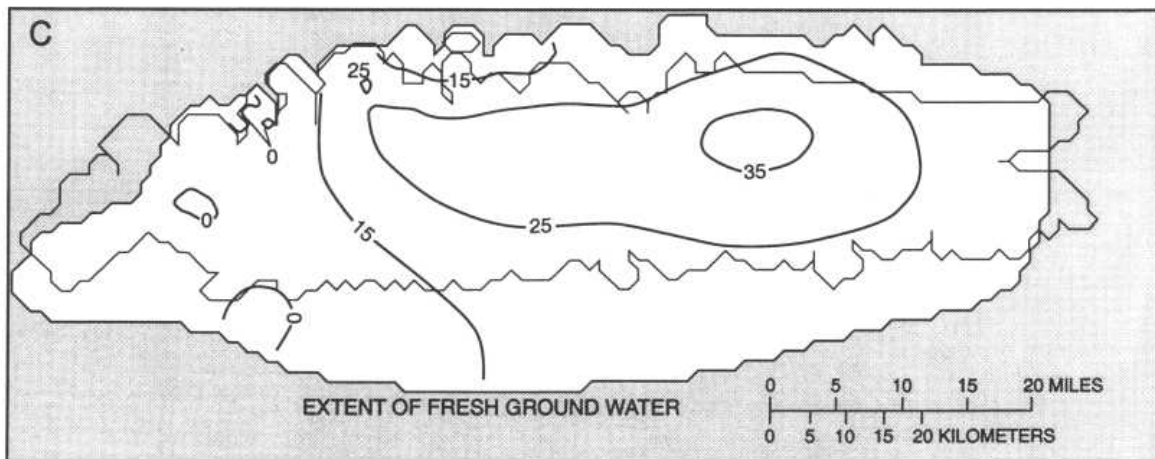
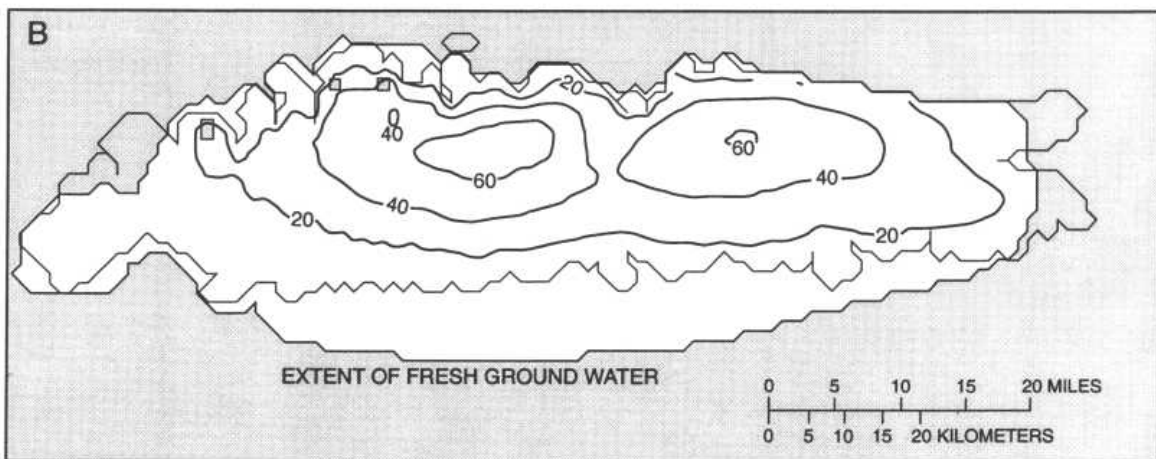
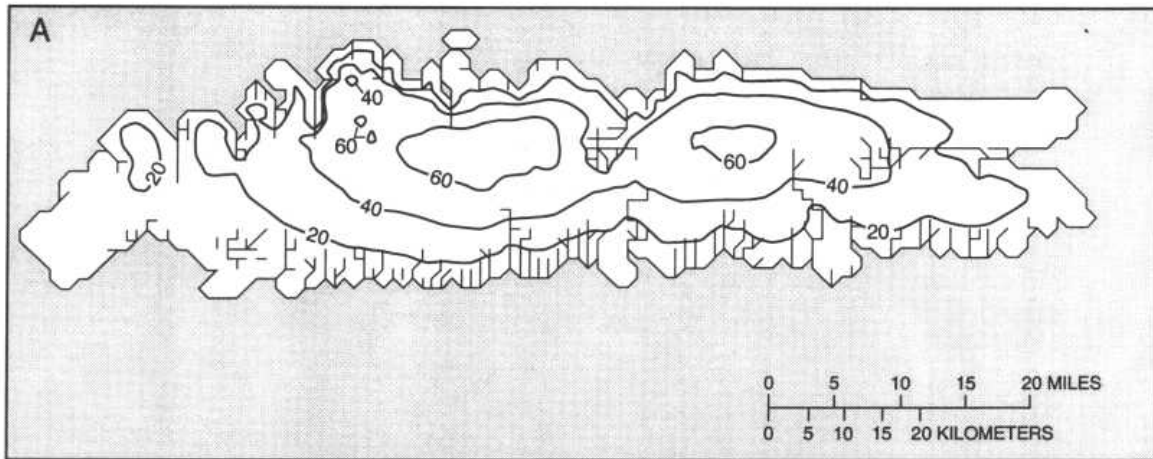
negligible however, because the average velocities would not cause movement of more than one model cell (4,000 ft). Although this assumption is acceptable at the regional scale and for estimating the ground-water flow budget and water levels, it does not address the possibility that saline ground water could be moving landward in local strata and in dilute concentrations sufficient to affect water supplies, particularly in southern Queens and Nassau Counties. Despite some water-level recovery in this area, water levels near the saltwater-freshwater interface in the Magothy and Lloyd aquifers have the greatest deficit in relation to the head needed to balance static sea water in these aquifers (fig. 31), indicating that this is the most likely place for saltwater

intrusion and, therefore, the best location for monitoring.

The interface generally is closer to shore on the northern shore than the southern shore. Small-scale pumping close to the northern shore could induce rapid local saltwater intrusion, particularly near the bays where erosion of confining units could have created a hydraulic pathway to confined aquifers.

Ground-Water Levels and Flow Patterns

The predicted distribution of hydraulic head in the three major aquifers in 2020 is shown in figure 31. (Corresponding maps for recent and predevelopment conditions are in figures 24 and 17, respectively.) The maps show that water levels in Kings, Queens, and



EXPLANATION

— 20 — POTENTIOMETRIC CONTOUR--Shows altitude at which water level would rise in a piezometer. Contour interval, in feet, is variable. Datum is sea level

Figure 31. Predicted ground-water levels for the year 2020. A., Water-table aquifer (model layer 1). B., Magothy aquifer (model layer 3). C., Lloyd aquifer (model layer 4).

most of western Nassau Counties in all three aquifers will recover from conditions during 1968-83 because pumping in Queens County will be decreased from 61 to 30 Mgal/d, and pumping in southwestern Nassau County will be replaced by pumping at the Muttontown Preserve pumping center. Stream lengths in this area also will increase (compare figures 31A and 24A).

The most severe water-level declines will be in central Nassau County, where the water table and potentiometric surface of the Magothy aquifer will decline 20 ft below levels during 1968-83 and 40 ft below predevelopment levels, on average. The potentiometric surface of the Lloyd aquifer in central Nassau will decline more than 10 ft below recent levels and 25 ft below predevelopment levels (fig. 31), and increased pumping from the Lloyd aquifer on the barrier island in southwestern Nassau County will result in increased drawdown very near the saltwater interface. Water-level declines in western Suffolk County will be smaller than in Nassau and will dissipate rapidly eastward; declines east of the Nissequogue and Connetquot Rivers will be only a few feet.

Although the total stress on the ground-water system in 2020 will be greater than that during recent conditions, the general redistribution of pumping away from severely affected areas of western Long Island will mitigate the severe drawdown below sea level in Queens and western Nassau Counties.

Ground-Water Budget

The ground-water budget for the year 2020 is shown in table 14; the values are derived solely from model-generated flow rates. The net stress on the ground-water system is 57 Mgal/d (24 percent) greater than during 1968-83 and represents an increase in pumping of 51 Mgal/d and a decrease in ground-water recharge from returned water of 6 Mgal/d. Under equilibrium conditions, the increase in

net stress is balanced by a corresponding decrease in discharge from the system. Comparison with the water budget for recent conditions (tables 9 and 14) indicates that 77 percent of the increased stress will be balanced by a net decrease in discharge to streams, and 12 and 11 percent by a net decrease in discharge to the shore and subsea boundaries, respectively. Comparison with the water budget for predevelopment conditions (tables 4 and 14) indicates that the total stress on the ground-water system in 2020 (297 Mgal/d) results in a decrease in base flow of 179 Mgal/d (39 percent), a decrease in discharge to shore boundaries of 89 Mgal/d (15 percent), and a decrease to subsea boundaries of 29 Mgal/d (36 percent). Islandwide, base flow is predicted to decrease to 86 percent of recent levels, or 61 percent of predevelopment levels.

Table 14. Ground-water budget for the year 2020 on Long Island

County	Recharge	Discharge			
	Precipitation and returned water use ¹	Pumpage ²	Stream	Shore	Subsea
Kings and Queens	136	46	22	66	3
Nassau	317	214	26	71	11
Western Suffolk	333	107	100	119	21
Eastern Suffolk	501	91	133	240	17
Total	1,287	458	281	496	52

¹Total recharge at the water table; includes water returned to the ground-water system after use and decreases due to increased runoff in Kings and Queens. (See table 11.)

²Includes total public-supply, industrial-commercial, and agricultural pumping.

The projected decrease in pumping in Kings and Queens Counties will cause increases in all components of natural discharge, including a 10-Mgal/d (83 percent) increase in base flow, a 10-Mgal/d (18 percent) increase in shore discharge, and a 1-Mgal/d (50 percent) increase in subsea discharge. Concern for increased flooding of underground struc-

tures or structures built near filled historic stream channels is warranted.

In Nassau County, a 58-Mgal/d increase in stress will decrease base flow to less than half of the amount under recent conditions and to about 20 percent of the predevelopment amounts (tables 4, 9 and 14). Shore and subsea discharge will decrease by 13 and 21 percent, respectively. Although the total decrease in subsea discharge warrants concern for salt-water intrusion, the increase occurs mostly in eastern Nassau, where the interface is far offshore and recent rates of intrusion are very slow.

In western Suffolk County, a 26-Mgal/d increase in stress will have considerably less effect. Base flow will decrease 19 percent from recent amounts, and shore and subsea discharge will decrease by 6 percent and 16 percent, respectively. In eastern Suffolk County, the increased stress will be only 4 Mgal/d, and the effects will be minor.

The increased pumping and sewerage in the 2020 water-supply strategy also will disturb the distribution of flow within the ground-water system (table 15). Recharge will decrease in Nassau and western Suffolk Counties because new sewerage will decrease the amount of returned water. Recharge will increase in eastern Suffolk, where pumpage will increase and cause a corresponding increase in returned water because most of the area is projected to be unsewered. Although the amount of water that flows deeper than the water-table aquifer will decrease in Kings and Queens Counties, significantly more water will flow to the deep aquifers on an islandwide basis. Ground water that flows to model layers deeper than layer 1 (generally below the water-table aquifer) increased from 462 Mgal/d under predevelopment conditions (table 5) to 648 Mgal/d during 1968-83 (table 10), and will increase to 681 Mgal/d by 2020 (table 15). Similarly, the amount of water that flows deeper than model layer 2 (generally the basal zone of the Magothy aquifer) increased from 235 Mgal/d

Table 15. Distribution of ground-water flow with depth for the year 2020 as represented in model

County	Model layer ¹			
	1 (water table)	2 (Magothy and Jameco)	3	4 (Lloyd aquifer)
Kings and Queens	136	42	31	5
Nassau	317	238	203	14
Western Suffolk	333	194	134	8
Eastern Suffolk	501	207	106	8
Total	1,287	681	474	35

¹Flow into layer 1 is recharge from precipitation and returned water use; flow into layers 2, 3, and 4 is leakage from the overlying layer.

under predevelopment conditions to 446 Mgal/d during 1968-83, and will increase to 474 Mgal/d by 2020. This information indicates that ground-water flow patterns, velocities, and residence times will be further altered by continual development. The most significant implication of which is that increased downward velocities to the deeper aquifers will increase the risk of contamination of those aquifers from land surface sources.

SUMMARY

Land use in Long Island ranges from highly urbanized and industrialized in the west to open land and agriculture in the east. In 1990-92, the population was nearly 6.9 million. Ground water is the sole source of water supply for Nassau, Suffolk and southeastern Queens Counties. In 1981, 385 Mgal/d was pumped for public supply, and an additional 115 Mgal/d was pumped for industrial-commercial and agricultural uses.

The Long Island ground-water system consists of a sequence of seven major hydrogeologic units. In order of deposition they are: the Lloyd aquifer, the Raritan confining unit, the Magothy aquifer, the Jameco aquifer, the Gardiners Clay (a confining unit), and the upper glacial aquifer. These units form a complex hydrogeologic framework that generally has

three major aquifer units whose degree of hydraulic connection varies locally, depending on the extent of intervening confining units.

This report describes the results of the simulation of the response of the Long Island ground-water system to water-supply and land development. Ground-water levels, base flow, and water budgets are provided for (1) predevelopment conditions (before-1900), (2) a severe drought in the 1960's, (3) conditions during 1968-83, and (4) the conditions that would likely result from a proposed water-supply development strategy for the year 2020. A three-dimensional ground-water flow model of the main Long Island ground-water system was used to provide quantitative estimates of these hydrologic conditions and of the relations between the hydrologic stress and the response of the ground-water system.

Before development, recharge from precipitation entered the Long Island ground-water system at an estimated rate of 1,126 Mgal/d; nearly 60 percent of which remained in the water-table aquifer; 37 percent moved to deeper units; and only about 3 percent entered the Lloyd aquifer. Recharge was balanced by discharge to streams (460 Mgal/d), the shore (585 Mgal/d), and subsea boundaries (81 Mgal/d). The water table attained a maximum altitude of more than 90 ft above sea level at the center of the island (near the Nassau-Suffolk County border) and contained prominent depressions near more than 100 ground-water-fed streams. The potentiometric surface of the Magothy aquifer was a subdued replica of the water table, and that of the Lloyd was considerably more subdued. The extensive Raritan confining unit severely retards ground-water flow to the Lloyd aquifer, but flow through local holes in this confining unit in Queens and northern Nassau Counties affects the source of water to and the shape of the potentiometric surface in the Lloyd aquifer.

Long Island's ground-water system is bounded laterally by saline ground water; the

saltwater-freshwater interface in the confined aquifers is offshore throughout most of southern Long Island, but lies close to the shore throughout northern Long Island. Ground-water levels in the confined aquifers indicate that the interface off the southern shore probably was moving landward, albeit slowly, even during the predevelopment period, in response to the sea-level rise since the last glacial period.

Development during the past 3 centuries has continuously affected the ground-water system of Long Island. Recharge from precipitation has been reduced by the paving of land surface, and large public-supply wells withdraw ground water from deep aquifers. Sewers discharge wastewater and in some places stormwater to the ocean, and stormwater infiltration basins augment recharge in Nassau and Suffolk Counties. Many streams in Kings and Queens have disappeared and subways and deep basements in parts of Kings now function as ground-water drains and require continuous dewatering.

By the early 1980's, more than 400 Mgal/d was pumped islandwide for public, industrial-commercial, and agricultural supplies, but some is returned as leakage from water-supply and sewer lines and as infiltration from domestic septic systems. Kings and Queens Counties import 700 Mgal/d from upstate reservoirs, and more than 50 Mgal/d of this probably reaches the ground-water system through leakage as an unintended form of artificial recharge.

The net stress on the ground-water system (reduced recharge and increased discharge as a consequence of development) during 1968-83 is estimated to be 240 Mgal/d. In response, base flow has decreased by 28 percent (135 Mgal/d). These effects are greatest in Kings, Queens, and western Nassau Counties, where water levels in all aquifers show considerable declines, and some cones of depression extend well below sea level. In these areas, the

saltwater-freshwater interface has moved landward, and low ground-water levels indicate that continued movement is likely. Monitoring between the interface and pumping centers would allow early detection of saltwater intrusion.

Simulation of the response of the ground-water system to the 1960's drought indicates that base flow of streams is sensitive to small water-table fluctuations, and that long streams are more sensitive than short ones. This is consistent with the observation that, during recent conditions, the reduction in base flow represents 56 percent of the net stress on the ground-water system.

A projected ground-water-supply strategy for the year 2020 was evaluated using the islandwide model. The net stress on the ground-water system was estimated to be 297 Mgal/d, an increase of 57 Mgal/d over 1968-83. The distribution of stress is expected to be dispersed over the island more uniformly than under recent conditions, however, and thus stress would decrease in Kings, Queens, and western Nassau Counties. As a result, ground-water levels would recover in western Long Island, mitigating the severe cones of depression and the landward gradients that threaten to induce saltwater intrusion in southwestern Long Island, and increasing the base flow in some streams. Most (77 percent) of the increased stress on the ground-water system would be balanced by decreased base flow, mainly in eastern Nassau and western Suffolk Counties; base flow would be reduced to about 20 percent of predevelopment levels in Nassau County and to about 70 percent in western Suffolk County.

The predicted ground-water system response to the proposed water-supply strategy for 2020 could best be used by comparison with the predicted effects of alternative strategies to identify the most effective methods to minimize the adverse hydrologic effects of development.

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